# DEVELOPMENT AND VALIDATION OF A REDUCED ONE-DIMENSIONAL THERMO-MECHANICAL SOIL STABILITY MODEL

For Predictive use with the Alaska RWIS

## **FINAL PROJECT REPORT**

by

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# SI\* (MODERN METRIC) CONVERSION FACTORS

C		MATE CONVERSIONS TO SI UNITS	
Symbol	When You Know	Multiply By To Find	Symbol
		LENGTH	
n	inches	25.4 millimeters	mm
t	feet	0.305 meters	m
yd	yards	0.914 meters	m
mi	miles	1.61 kilometers	km
		AREA	
in <sup>2</sup>	square inches	645.2 square millimeters	mm <sup>2</sup>
ft <sup>2</sup>	square feet	0.093 square meters	m <sup>2</sup>
yd <sup>2</sup>	square yard	0.836 square meters	m <sup>2</sup>
ac	acres	0.405 hectares	ha
mi <sup>2</sup>	square miles	2.59 square kilometers	km²
		VOLUME	
fl oz	fluid ounces	29.57 milliliters	mL
	gallons	3.785 liters	L
gal ft³	cubic feet	0.028 cubic meters	m <sup>3</sup>
yd <sup>3</sup>	cubic yards	0.765 cubic meters	m <sup>3</sup>
,~		umes greater than 1000 L shall be shown in m <sup>3</sup>	
		MASS	
oz	ounces	28.35 grams	a
lb	pounds	0.454 kilograms	g kg
		or roll	
Т	short tons (2000 lb)	0.907 megagrams (or "metric ton")	Mg (or "t")
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°F	Fahrenheit	5 (F-32)/9 Celsius	°C
		or (F-32)/1.8	
		ILLUMINATION	
fc	foot-candles	10.76 lux	lx
fl	foot-Lamberts	3.426 candela/m <sup>2</sup>	cd/m <sup>2</sup>
	FOR	RCE and PRESSURE or STRESS	
lbf	poundforce	4.45 newtons	N
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#### **EXECUTIVE SUMMARY**

Seasonal vehicular weight road restrictions are often used in Northern climates to mitigate the damage that can occur in spring during thaw. Frost heave and related freeze/thaw processes in wet soil decrease the compressive strength of highways. The location of the frozen/unfrozen interface is a primary determinant for the physical stability of the overlying material. The Alaska Dept. of Transportation and Public Facilities (ADOT&PF) generally apply weight restrictions during spring thaw until the thaw depth reaches five feet. ADOT&PF maintains a system of Road Weather Information Stations (RWIS) that record air and subsurface temperatures in order to know approximately when this occurs. There are some limitations to relying on RWIS sites for determination of seasonal commercial vehicle weight restrictions, such as cost and the associated limited number of sites over a large geographic area. A numerical model was developed as a one-dimensional finite difference thermal energy balance that accounts for both sensible and latent heat effects in a semi-saturated soil. The primary goal was to accurately forecast the time when specific vertical locations thaw; specifically, 1 foot and 5 feet below surface in this case. The tool is in the form of an Excel spreadsheet and requires only a single adjustable parameter that represents a modified thermal diffusivity that accounts for latent heat effects. The model is driven using air temperature data, and predicts the temperature at depth. The time resolution of the model is the same as the input temperature data. The model was calibrated and evaluated using archived RWIS data at five locations in Alaska with 1 hour resolution. The same value for apparent thermal diffusivity was used at all sites. The average error at 1 foot depth was 4.78 days, and 8.5 days at 5 feet depth. The spreadsheet tool does not include any macros or external scripts, so can be used in most other major spreadsheet programs such as Libreoffice Calc.

### CHAPTER 1. INTRODUCTION

Seasonal vehicular weight road restrictions are often used in Northern climates to mitigate the damage that can occur in spring during thaw. Frost heave and related freeze/thaw processes in wet soil decrease the compressive strength of highways. Frozen ground contains ice in different forms ranging micronscale soil particle coatings to millimeter-scale ice inclusions. When thawing, the associated volume-contraction of water leads to a decrease in mechanical strength of the highway. The overburden pressure of vehicle traffic causes the soil skeleton to adapt. Concurrently, the excess pore pressure eventually decreases as liquid drains; and the rate is a function of the solid material physical characteristics and the thermal regime.

The location of the frozen/unfrozen interface is a primary determinant for the physical stability of the overlying material. Its relative impermeability severely impedes relaxation of the excess pore water pressure. Therefore, depth-of-thaw location is conventionally used as an indicator for a highway's seasonally changing mechanical strength. The thaw depth can be measured automatically in near-real time using temperature depth probes (TDP), which are most often a string of thermistors embedded in a rigid rod at strategically determined depth locations along the probe rod.

The Alaska Dept. of Transportation and Public Facilities (ADOT&PF) maintains a system of Road Weather Information Stations (RWIS) along the primary highway system in Alaska. Each station includes an Environmental Sensor Station (ESS) that measures and records atmospheric, surface/sub-surface, and water/snow conditions at the site. A TDP is used to measure the depth of thaw. The TDPs in the Alaska RWIS system have thermistors located at 3-inch intervals for the first foot below surface, and then 6-inch intervals for an additional five feet of depth. Data collected by the ESS are collected by a remote processing unit at the RWIS, and then transmitted every 6 hours via telephone (where available) or an alternative wireless communication system. A subset of the collected ESS data is then made available in graphical and tabular form on the web through the main AKDOT&PF portal.

There are some limitations and weaknesses in the current use of TDP measurements at RWIS sites for determination of highway material strength and therefore seasonal commercial vehicle weight restrictions. One limitation is the finite and few number of RWIS locations throughout the geographically large Alaska highway system. Clearly the financial cost of installing and maintaining each station limits the number and density of sites. Another limitation is the finite number and spacing between thermistor sensors, which results in a somewhat low-resolution measurement of the actual frost depth. Finally, equipment failure of TDPs results is not uncommon, yet difficult to repair timely if at all. There are also weaknesses in using only depth of thaw as the determining factor for weight restrictions. Various subgrade aggregate materials and natural soils both drain at different rates, as well as have differing cohesive strengths as functions of water content and pore pressure.

### 1.1. Road Materials

The soil used under highways is typically carefully selected and engineered to provide a stable foundation for the road structure. This subgrade is typically native soil upon which the road is built. It's often compacted and sometimes treated to improve its properties. Ideal subgrade soils are well-draining and have good bearing capacity.

Above the subgrade is the sub-base. This is typically gravel, crushed stone, and/or sand. The sub-base helps with drainage and provides a stable platform for the base course. In areas where the native soil is unsuitable, engineered fill may be used. This is specially selected or treated soil that meets specific engineering requirements.

Key characteristics of soils used under highways include good drainage properties, high bearing capacity, low frost susceptibility, minimal shrink-swell potential, and compactability to achieve desired density. The exact type and composition of soil used can vary based on local availability, climate conditions, traffic load expectations, and regional engineering practices.

## 1.2. RWIS System

The RWIS system, or Road Weather Information System, is a network of environmental sensor stations used to collect and disseminate road weather data. It's a crucial tool for highway maintenance and safety, especially in regions with challenging weather conditions. RWIS stations typically collect air temperature, humidity, wind speed and direction, precipitation type and intensity, pavement temperature, subsurface temperatures, and sometime pavement condition (e.g. wet, dry, icy). RWIS systems play a crucial role in managing roads in areas prone to freezing, helping to prevent accidents and optimize maintenance efforts. They provide real-time and forecast information that is valuable for both immediate operational decisions and long-term planning. RWIS provides crucial data that helps road maintenance teams make informed decisions about how to manage these conditions.

## 1.3. Project Goals

The numerical model developed in this project is a one-dimensional finite difference thermal energy balance that accounts for both sensible and latent heat effects in a semi-saturated soil. The primary goal was to accurately forecast the time when specific vertical locations thaw; specifically, 1 foot and 5 feet below surface in this case.

The number of initial input parameters into the model were kept as small as possible. Initially the set would include the physical properties of the underlying soil materials (thermal conductivity, heat capacity, and density). Effective freezing and thawing indices (n-factors) will be determined using archived RWIS data of temperature profiles over time. Effective surface temperatures can then be matched to air temperatures for a top-surface boundary condition into the one-dimensional model.

Initial validation of the model involved comparison of the model forecasts with several years of archived Alaska RWIS data from TDP measurement sites. We worked with Alaska DOT&PF for acquisition of the data, and to gain an understanding of the conditions at each of the RWIS locations. The model validation metrics were deviation between forecast and recorded temperatures with focus on temperatures near freeze/thaw. Model design iteration included reducing the model input requirements such that it can be modified and implemented efficiently for use in a wide range of locations. The ulterior motive of this step is to be applicable to locations far from any current RWIS installations. After the development, validation, and iteration steps, the forecast system wasdesigned for facile integration with the current Alaska RWIS data system.

#### CHAPTER 2. LITERATURE REVIEW

One-dimensional finite difference models (FDMs) have become a cornerstone for simulating the freezing process in soils. FDMs have been widely employed to simulate one-dimensional freezing processes in soils due to their relative simplicity and computational efficiency. This review explores the strengths and limitations of FDMs in this context, highlighting key considerations for model development and application. This review examines key developments and applications of these models in geotechnical engineering and permafrost studies.

Heat transfer in soil freezing is governed by the heat equation, incorporating latent heat release during water-to-ice phase change. FDMs discretize the spatial domain (usually depth) and time domain, approximating the temperature distribution within the soil profile. Early models often assumed a constant soil composition and a single freezing point for all water (Jumikis, 1977). However, advancements incorporated the dependence of unfrozen water content and thermal properties on temperature (Kozlowski, 2001; Riseborough & Smith, 1985). This improved the accuracy of frost depth prediction, a crucial parameter in geotechnical engineering.

The complexity of FDMs for soil freezing can vary. Simpler models may focus solely on heat transfer, while more sophisticated approaches couple heat and mass transfer to account for water movement during freezing (Comes-Pintaux & Nguyen-Lamba, 1986). The selection depends on the specific problem being addressed and the desired level of detail.

Several factors influence the accuracy and applicability of FDMs. Mesh resolution (spatial and temporal discretization) is critical, as finer meshes lead to more accurate results but require greater computational resources. Furthermore, reliable soil property data, including thermal conductivity, volumetric heat capacity, and the relationship between unfrozen water content and temperature, are essential for accurate simulations (Goodrich, 1978) (Harlan, 1973).

Validation of FDM simulations is crucial. Comparisons with analytical solutions (for simple cases) or field measurements provide confidence in the model's performance. FDMs have been successfully applied to various problems, including predicting frost depth in foundations, evaluating the effectiveness of ground freezing techniques, and understanding the impact of climate change on permafrost. Early work by Harlan (1973) established a foundation for modeling coupled heat and mass transfer in freezing soils using finite difference methods. His model incorporated both conductive and convective heat transfer, as well as moisture migration driven by temperature gradients. Nixon (1975) considered the role of convective heat transport in the thawing of frozen soil. He concluded that for a wide range of conditions, the effect is minor and does not play a significant role in determining the rate of thaw.

Subsequent research focused on improving the representation of soil properties and phase change dynamics. Jame & Norum (1980) developed a model that accounted for the variation of unfrozen water content with temperature, a crucial factor in accurately simulating frost heave and thaw settlement. Konrad & Morgenstern (1984) introduced the segregation potential concept to model frost heave, which was later incorporated into finite difference schemes by various researchers. This approach allowed for more accurate prediction of frost heave in frost-susceptible soils.

Recent advancements have focused on incorporating more complex phenomena and improving numerical stability. Hansson, et al. (2004) developed a model that accounts for salt transport and its

effects on freezing point depression, which is particularly relevant for coastal and saline soils. Dall-Amico, et al. (2011) presented a robust numerical scheme for solving the coupled heat and water flow equations in freezing soils, addressing issues of convergence and stability in previous models. Current research trends include the integration of finite difference models with other numerical techniques, such as finite element methods, to handle more complex geometries and multidimensional problems.

### 2.1. Latent Heat of Fusion

The latent heat of fusion is a critical component in finite difference models of soil freezing, as it represents the energy required for phase change between water and ice. The apparent heat capacity method is one of the most common approaches. The latent heat is incorporated into an "apparent" or "effective" heat capacity of the soil-water-ice system. The heat capacity is treated as a function of temperature, with a large spike around the freezing point to represent the latent heat effect. This method, used by researchers like Hansson et al. (2004), allows for a smooth transition between frozen and unfrozen states.

Some models, like those based on Harlan's (1973) work, treat the latent heat as a source or sink term in the heat transfer equation. As freezing occurs, the latent heat is released (acting as a heat source), and during thawing, it's absorbed (acting as a heat sink).

This enthalpy formulation, employed by Dall'Amico et al. (2011), uses enthalpy as the primary variable instead of temperature. The enthalpy includes both sensible and latent heat, allowing for a more natural incorporation of phase change effects.

In models that explicitly track the freeze-thaw interface, like Nixon's (1975), the latent heat is accounted for as a boundary condition at the moving frost front. This approach is particularly useful for sharp freeze-thaw interfaces. Many models incorporate the relationship between unfrozen water content and temperature. This indirectly accounts for latent heat effects, as the gradual release of latent heat is reflected in the changing unfrozen water content over a range of sub-zero temperatures.

The choice of method often depends on the specific problem, soil characteristics, and desired balance between computational efficiency and accuracy. More recent models tend to use combinations of these approaches to better represent the complex thermodynamics of freezing soils.

#### CHAPTER 3. DATA AND METHODS

Determining the physical properties of soil is crucial for accurate finite difference modeling of freezing processes. Models typically handle soil properties in a variety of ways. Many models rely on experimentally determined soil properties as input parameters. These may include thermal conductivity, heat capacity, hydraulic conductivity, porosity, bulk density and particle size distribution. Researchers often conduct laboratory tests on soil samples to obtain these properties before running simulations.

Some properties, particularly thermal properties, are often represented as functions of temperature and phase composition. For example, thermal conductivity is often modeled using weighted averages of soil components (solid particles, water, ice, air) based on volume fractions, which change with temperature. Heat capacity can be similarly calculated using mixture models, accounting for changing proportions of water and ice.

The critical relationship between temperature and unfrozen water content is typically determined experimentally for specific soil types. It's often represented by empirical functions (i.e. unfrozen water curve) or lookup tables in the models.

Some models use parameterization schemes to estimate properties based on soil classification (e.g., sand, silt, clay percentages) and other basic soil data. Often, initial estimates of soil properties are refined through model calibration against field or laboratory data, improving the accuracy of simulations.

It's worth noting that the accuracy of many models heavily depends on the quality and appropriateness of the soil property inputs. Sensitivity analyses are often performed to understand how uncertainties in soil properties affect model outcomes. It was beyond the scope of this project to obtain physical samples and measure physical properties at the sites being studied. Furthermore, it aligns with the overall project goal to keep input parameters to a minimum. Therefore, the initial approach, which was also the final product, required only the thermal diffusivity of the material

$$\alpha = \frac{k}{\rho \ c_p}$$

Furthermore, this value remained the same at all sites investigated. This was not unexpected since the material physical properties down to 5 feet are very similar at all locations.

## 3.1. Numerical Methods

This model solves the unsteady, one-dimensional heat conduction equation

$$\frac{\partial T}{\partial t} = \alpha \frac{\partial^2 T}{\partial z^2}$$

by the finite difference method. The latent heat of fusion accounted for using the apparent heat capacity method since tracking the location of the freezing interface would be very difficult with the course resolution and fixed grid of this model.

The entire model is contained in a single Microsoft Excel spreadsheet and contains no macros or other external functions and methods. The only required temperature input into the model is the RWIS

measured air temperature as a function of time. Using that value, the temperature at the following depths (in feet) are calculated: 0.25, 0.5. 0.75, 1.0,1.5, 2.0, 2.5, 3.0, 3.5, 4.0, 4.5, 5.0, 5.5.

The upper most depths were able to be adequately determined without needing to account for latent heat effects. Down to the 1.0 feet depth, the temperature can be determined by an analytic solution of one-dimensional, time-dependent conduction equation with constant physical properties.

$$\frac{T(z,t) - T_i}{T_s - T_i} = erfc\left(\frac{z}{2\sqrt{\alpha t}}\right)$$

where  $T_s$  is the temperate at the cell directly above it, and  $T_i$  is the temperature of the same cell in the time step before current. Below this depth, the thermal diffusivity is determined by a conditional test using the temperature relative to the freezing point.

#### 3.2. RWIS Data

Archived surface and sub-surface temperature data were acquired through the Alaska Department of Transportation and Public Facilities Road Weather Information System website at https://roadweather.alaska.gov/gis.. For all five of the sites used in this study, six years of historical temperature data was acquired. This data included air temperature and ground temperature at all the depths given above. The input to the model is the air temperature, and the remaining ground temperatures were used for development, adjustment and tuning of the model.

## 3.3. Thermal Model Spreadsheet Tool

This section is a brief visual overview of what the thermal model spreadsheet tool looks like to the end user. Note that the actual sheet may have up to 10,000 columns depending on the time resolution and duration of the data. Time increments of 1.0 hours were used in the development of this tool. Different time increments can be entered for the air temperature, however, and it can be either regularly or irregularly spaced.

Only the left-most part of a single tab of the spreadsheet is usually shown in most of the figures below. Figure 1 shows an overview of what the Excel spreadsheet looks to the user. One visual effect added to the cell data background was color coding from blue to red (cold to warm). For model validation, the actual RWIS data at all depths is included in the data block below the model forecasts. The top block of cell is the model forecast, and the block below that is the actual measure temperature. The top scatter plot is the temperature at 1.0 feet showing both the model and actual measurement, and the plot below that is the same for 5.0 feet depth.

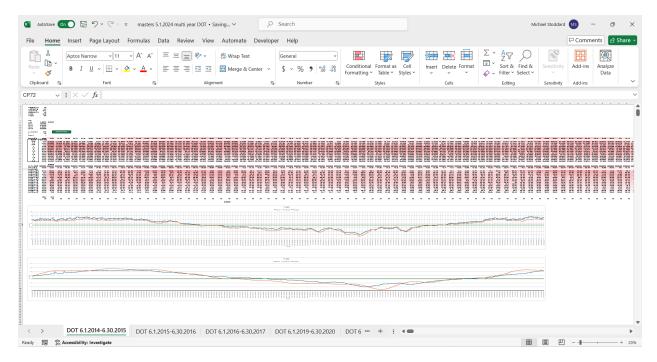


Figure 1 Overview of the model interface

Figure 2 shown below is a closer view of the model data block using background color coding to aid in quick visualization of where the freezing interface is located. Red colors are above freezing, and blue colors are below freezing. Each column represents a time difference of 1.0 hours, and the rows are at the depths given earlier (most are 0.5-foot increments). In this figure, a general idea of the descending freezing front can be seen in the white band.

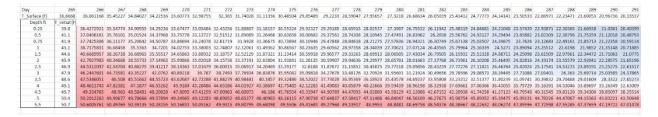


Figure 2 Closer view of the data blocks with background color coding

Figure 3 simply illustrates a user interfacing with the model. User adjustable parameters are in the upper left. There are a few cells below those that were used for development and tuning of the model and can be hidden or even completely removed by the user if desired. Time data are entered in a single row of the model in any resolution available to or desired by the user.

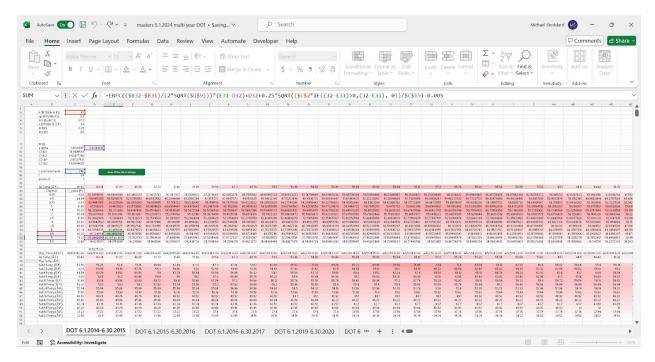


Figure 3 Illustration of a user interfacing with the model

Figure 4 below is a close-up section view of the temperature plotted as a function of time at the 1.0 depth (top) and 5-foot depth (bottom). The orange line is the simulated at-depth temperature, and the blue line is the actual temperature recorded from RWIS. The straight green line at 32 degrees Fahrenheit shows the bulk freezing temperature. The road material at 1-foot and 5-foot depth does not contain any appreciable unfrozen water below that temperature, and there is a negligible freezing point depression. Therefore, the green line is a fairly accurate location of the interface between unfrozen and frozen material. The red arrows nearer the right side of each plot show the predicted and actual time of thaw, which are nearly identical in this simulation.

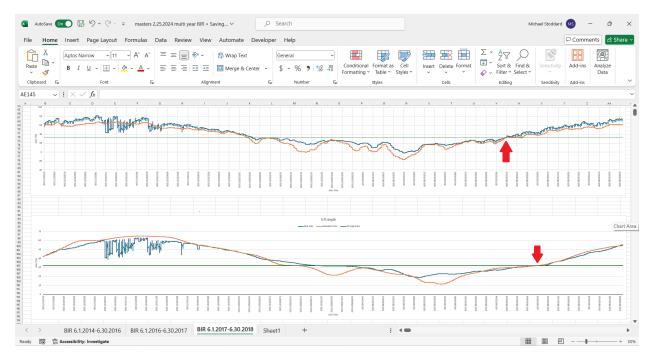


Figure 4 Scatter plots of the 1 foot and 5 foot depths, including the actual and model predictions

### CHAPTER 4. FINDINGS

The model was calibrated and applied at four Alaska locations. In order to fully evaluate the predictive capability of the model, sufficient archived air, surface and subsurface temperature data are necessary Therefore, some sites (e.g. DOT Lake) had more comparison than others (e.g. Nenana Hills).

Each location was put into a separate copy of the master spreadsheet with tabs typically in one-year intervals due to the size constraint of Excel columns. The four sites used in this analysis are shows in the figure below:

- NEN- Parks Highway @ Nenana Hills MP 325.4
- DOT-Alaska Highway @ Dot Lake MP 1355.2
- CLR- Steese Highway @ Cleary Summit MP 20.9
- BIR- Richardson Highway @ Birch Lake MP 307.2

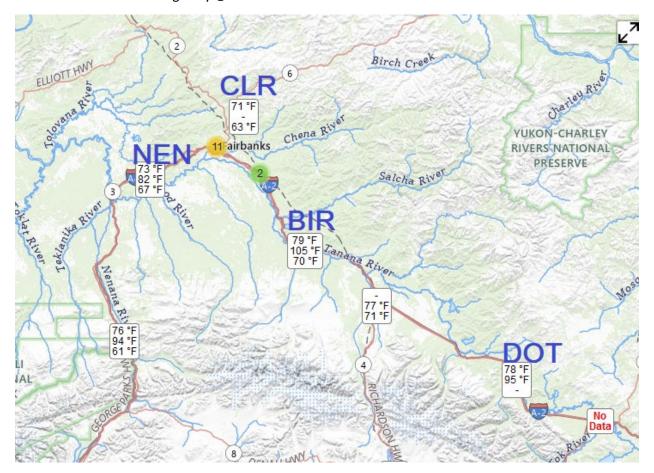


Figure 5 Map of the Interior of Alaska showing the four FWIS sites used in this analysis

Although the model predicts temperatures at all the depths discussed in section 4.1, the primary focus was on thaw at depths of 1.0 and 5.0 feet. Table 1 below summarizes the model performance for each of the sites analyzed.

Table 1 Summary of the predicted and actual dates of thaw for each site analyzed

Site	1 ft (model)	1ft (actual)	Error [d]	5 ft (model)	5 ft (actual)	Error [d]
NEN	4/12/18 17:00	4/12/18 22:00	0.21	5/6/18 5:00	5/12/18 20:00	6.63
DOT	4/8/15 19:00	4/14/15 18:00	5.96	5/6/15 11:00	5/16/15 3:00	9.67
	3/28/16 8:00	3/31/16 7:00	2.96	4/21/16 5:00	5/5/16 18:00	14.54
	4/15/20 21:00	4/20/20 19:00	4.92	5/6/20 1:00	5/14/20 6:00	8.21
	4/19/21 16:00	4/20/21 17:00	1.04	5/13/21 19:00	5/12/21 1:00	1.75
	4/25/22 10:00	4/25/22 8:00	0.08	5/17/22 4:00	5/21/22 5:00	4.04
BIR	3/27/15 13:00	4/3/15 22:00	7.375	N/A	N/A	N/A
	3/29/16 16:00	3/28/16 21:00	0.792	4/22/16 22:00	4/23/16 9:00	0.458
	4/5/17 20:00	4/5/17 20:00	0	4/30/17 18:00	5/3/17 9:00	2.63
	4/11/18 18:00	4/11/18 17:00	0.042	5/3/18 14:00	5/6/18 20:00	3.25
CLR	3/28/15 17:00	4/19/15 20:00	22.13	7/4/15 1:00	N/A	N/A
	4/6/16 15:00	4/11/16 21:00	5.25	4/22/16 15:00	5/13/16 14:00	20.96
	4/12/17 11:00	4/23/17 1:00	10.58	5/7/17 16:00	5/20/17 22:00	13.25
	4/26/18 10:00	4/23/18 22:00	2.5	5/14/18 21:00	5/25/18 20:00	10.96
	4/15/20 22:00	4/23/20 20:00	7.92	5/11/20 9:00	5/25/20 12:00	14.13

There is a row in the table for each year that there was sufficient data to do a comparison the model performance. Occasionally there were data gaps around the time of thaw that precluded a complete comparison; those are indicated with N/A in the table. The model time step used was one hour since that corresponds with the RWIS data resolution. Therefore, the date and time of thaw for both the model and the actual have an uncertainty of about 1 hour. As discussed in Section 4.3, the model can accommodate any time step and depends only on the resolution of the air temperature driving the simulation. The absolute difference between the model predictions and the actual thaw are shown in the Error column in units of days.

A summary of basic statistics of the error was calculated using all sites and all years, weighing each instance equally. A summary of these statistics is shown in Table 2 below.

Table 2 Statistic summary of the error

	1	5
	foot	foot
	error	error
	[d]	[d]
Minimum	0	0.46
Maximum	22.13	20.96
Average	4.78	8.50

Predictions for the Cleary Summit (CLR) on the Steese highway had substantially more error than the other four sites; approximately double for all five years analyzed. Since CLR had one of the most comprehensive data sets (five years), the discrepancies of that site alone weigh more heavily in the statistics. Insets showing time of thaw at CLR are shown in Figure 5 below. It appears there is a larger

latent heat effect at this location and an adjustment of the thermal diffusivity would yield closer agreement, but the investigation is continuing as part of the larger graduate student's research project.

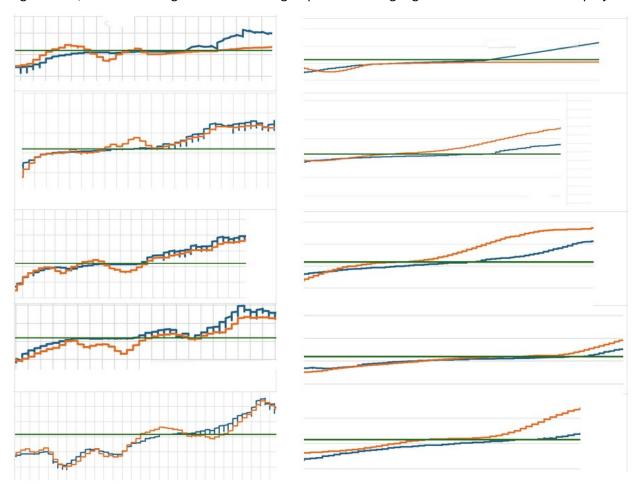


Figure 6 Model (yellow) and actual (blue) during thaw at 1 and 5 feet

### CHAPTER 5. CONCLUSIONS

A numerical tool was developed that can fairly accurately predict the time of thaw at depth below a roadway using only surface air temperature data. The tool comes as an Excel spreadsheet and contains no macros or other external programs/scripts, so it can easily be used in many other major spreadsheet programs such as the free open-source LibreOffice Calc without modification. The uncertainty in the model is the same as the time resolution of the air temperature data driving the model. There is a single adjustable parameter that represents an effective thermal diffusivity that accounts for latent heat effects. A single value for this parameter was used in the analysis of CHAPTER 5, but an end user can easily adjust this value for a different site if there is some archived data available to validate it. We found a single value worked fairly well for the five sites analyzed because the subsurface material is likely similar at all sites. The average error at a depth of 1 foot was 4.78 days, and the average error at 5 feet was 8.5 days.

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